

# Development of 3D-Coupled hydrodynamics-ecosystem model for Jakarta Bay

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## Introduction

Jakarta bay, as one of the entering gate to the Jakarta City, is one of the important waters in Indonesia. According to this function, Jakarta bay has too many changing in hydrodynamics conditions caused by nature and human activities. The study of current circulation in Jakarta bay was very important. To study the ocean circulation and physical processes in Jakarta Bay, a numerical modeling system based on the Princeton Ocean Model (POM Model, Blumberg and Mellor, 1987) has been implemented in the Jakarta Bay.

Although it is clear that Jakarta Bay is heavily impacted by human activities, it is not clear how the organic inputs interact with physical conditions to influence the lower trophic level ecosystem. In particular, the physical processes that control the ecosystem are not well understood, so here we examine the interaction between the physical and biological processes that drive the temporal variation and spatial distribution of nutrients and phytoplankton in Jakarta Bay, using a coupled hydrodynamic–ecosystem model.

A joint collaboration research between Laboratory of Marine Survey Technology (BTSK)-BPPT and Center for Marine and Environmental Studies (CMES) have been collected the observation data such as sealevel, temperature-salinity and dissolved oxygen (DO) profiling, ocean current observation, nutrient sampling, etc since December 2015-December 2018. These observation data used to perform the above ocean model in Jakarta Bay.

In this study, we only consider the tidal and wind monsoon as principal forcing to generate the physical processes in Jakarta Bay waters.

## Aims

The aim of this study are, i.e:

- to understand the current circulation and to predict the physical processes in Jakarta Bay using the POM model.
- to predict the distribution of ecosystem parameter such as Dissolved Inorganic Nitrogen (DIN), Dissolved Inorganic Phosphate (DIP), phytoplankton type-1 (PHY), zooplankton (ZOO), Detritus (PON) and Dissolved Oxygen (DO)

## Procedures

### Numerical Modelling

The Princeton Ocean Model (POM) is a 3-D primitive equations ocean circulation model, which has been described in detail by Blumberg and Mellor (1987) and Mellor (2004). POM is a free surface, sigma coordinate model that uses a time-splitting technique to solve depth-integrated and fully three-dimensional equations with different time steps. The surface elevation, velocity, temperature, and salinity fields are prognosticated assuming the fundamental hypotheses: (i) that the seawater is incompressible, ii) that the pressure in any point of the ocean is equal to the weight of the column of water over it (hydrostatic approximation), and iii) that the density can be expressed in terms of a mean value and a small fluctuation (Boussinesq's approximation). Consequently, in a system of orthogonal Cartesian coordinates, the governing equations are:

$$\frac{\partial U_i}{\partial x_i} = 0 \quad (1)$$

$$\frac{\partial(U,V)}{\partial t} + \frac{\partial[U_i(U,V)]}{\partial x_i} + f(-V, U) = \rho_0^{-1} \left[ \frac{\partial p}{\partial x}, \frac{\partial p}{\partial y} \right] + \frac{\partial(K_M \frac{\partial(U,V)}{\partial z})}{\partial z} + (F_U, F_V) \quad (2)$$

$$\frac{\partial p}{\partial z} = -\rho g \quad (3)$$

$$\frac{\partial \theta}{\partial t} + \frac{\partial(U_i \theta)}{\partial x_i} = \frac{\partial(K_H \frac{\partial \theta}{\partial z})}{\partial z} + F_\theta + \frac{\partial R}{\partial z} \quad (4)$$

$$\frac{\partial S}{\partial t} + \frac{\partial(U_i S)}{\partial x_i} = \frac{\partial(K_H \frac{\partial S}{\partial z})}{\partial z} + F_S \quad (5)$$

Which are  $K_M$  and  $K_H$ , are computed according to the Mellor-Yamada 2.5 turbulence closure scheme (Mellor and Yamada, 1982). This scheme requires prognostic calculations of  $(q^2)$  and  $(q^2 l)$ , where the turbulent kinetic energy is  $q^2/2$  and the turbulent macro scale is  $(l)$ . The equations that calculate  $q^2$  and  $q^2 l$  are similar to those used for temperature and salinity and include production and dissipation source/sink terms. Horizontal diffusion terms ( $F_U, F_V$ ,  $F_\theta$ , and  $F_S$ ) are computed using the Smagorinsky (1963) formulation, implemented into POM according to Mellor and Blumberg (1986). The last term in Eq. (4), is the solar radiation flux that penetrates the sea surface. The UNESCO equation of state, as adapted by Mellor (1991) is used. The in situ density is determined as a function of salinity, potential temperature and pressure; the latter being approximated by the hydrostatic relation and constant density.

### Model Configuration

The model domain comprises the Jakarta Bay, Indonesia, from 106° 33' 00"-107° 01' 30"E and from 6 ° 12' 00" – 5° 49' 30"S. The bottom topography used the Shuttle Radar Topography Mission (SRTM) with 30 second grids ([https://topex.ucsd.edu/WWW\\_html/srtm30\\_plus.html](https://topex.ucsd.edu/WWW_html/srtm30_plus.html)). Grid resolution model constructed from the topogarchy data for 1/400° x 1/400° degree. In this model version (afterward JB model), there are three open lateral boundaries; at 106.555 °E, at 107.025 °E and 5.825 °N. The JB model has been configured with a horizontal resolution of 1/400° in latitude by 1/400° in longitude. In the vertical , the grid has 11 sigma layers with logarithmic distributions at the surface and the bottom and a linear distribution in between. Other significant model parameters are given in **Table-1**.

**Table-1.** Model Parameter used in The Jakarta Bay (JB) model.

Parameter	Description Value
C	Non-dimensional constant used in calculating 0.2
$\nu$	the horizontal viscosity for momentum
$\tau_{prni}$	Inverse horizontal Prandtl Number 0.2
$\mu M$	Background viscosity 0.00002
$Z_o$	Bottom roughness length 0.01
$\Delta t_{ext}$	External mode time step 3 sec
$\Delta t_{int}$	Internal mode time step 60 sec

Open lateral boundaries conditions which are derived from FES2016 Global Tide Model are modified, e.g. tidal amplitude of FES2016 + 0.008m. Therefore, open lateral boundary of barotropic tidal currents which are derived 2-D tidal model are applied to the 3-D tidal model. Surface boundary applied in this model are only the wind data collected from Meteorological Maritime Station at Tanjung Priok Port, Jakarta (106° 52' 49.7"E, 06° 06' 28.3" S). Time simulation period for the model is started from June 1, 2015 untill December 31, 2016 (1.5 years).

### **Ecosystem Model**

The ecosystem model (Figure 2) has five compartments, namely nitrate (NO<sub>3</sub>), ammonium (NH<sub>4</sub>), phytoplankton (F), zooplankton (Z), and benthic detritus.

## Results

In the following paragraph, the harmonic analyses for four constituents (M2,S2, K1 and O1) from model results compare to the tide observations as reported by Koropitan and Ikeda (2008) on a tidal station at in Jakarta Port (106.89083°E, 6.10667°S). Table-2 show the model is very close to the observation. Some small differences of phase is considered as mixing process of water masses in the Jakarta Bay.

After the wind forcing considered in the JB's 3D-Tide model, we did the monthly-mean of surface horizontal distributions of sea level and current velocity fields from the daily model results. The analysis started from August 2015 to December 2016 due to neglect the spin-up time period (~about 2 months).

In August-September 2015 the vector fields flows from the east to west side of Jakarta Bay. Vector field on September 2015 tends to increase due to the moonsoon winds change from east to east-west transition monsoon winds. Sea surface tends to increase on western side of Jakarta Bay on August-September 2015. But, starting in October-December 2015 the increasing of sea surface propagates to the center and eastern side of Jakarta Bay (**Figure-1**).

**Table-2.** Comparison of observed and modelled M2,S2,K1 & O1 tidal elevations at the coastal tide-gauge station at Jakarta (106° 52' 49.7"E, 06° 06' 28.3" S).

Constituents	Model	Observed*
Amplitude (cm)		
Phase (oG)		
M2	5.41	5.45
	140.85	134.06
S2	5.04	5.08
	78.08	73.69
K1	25.17	24.92
	34.73	32.69
O1	13.75	13.70
	25.32	22.84

*\*Koropitan and Ikeda,2008*

Again, the surface currents tends to flow from west to east side Jakarta Bay during December 2015 to February 2016. But in this period, there are three positions of high sealevel in Jakarta Bay, i.e. east side (December 2015), west side (January 2016) and east side (February 2016). It means the model results shows the anomaly sea surface in

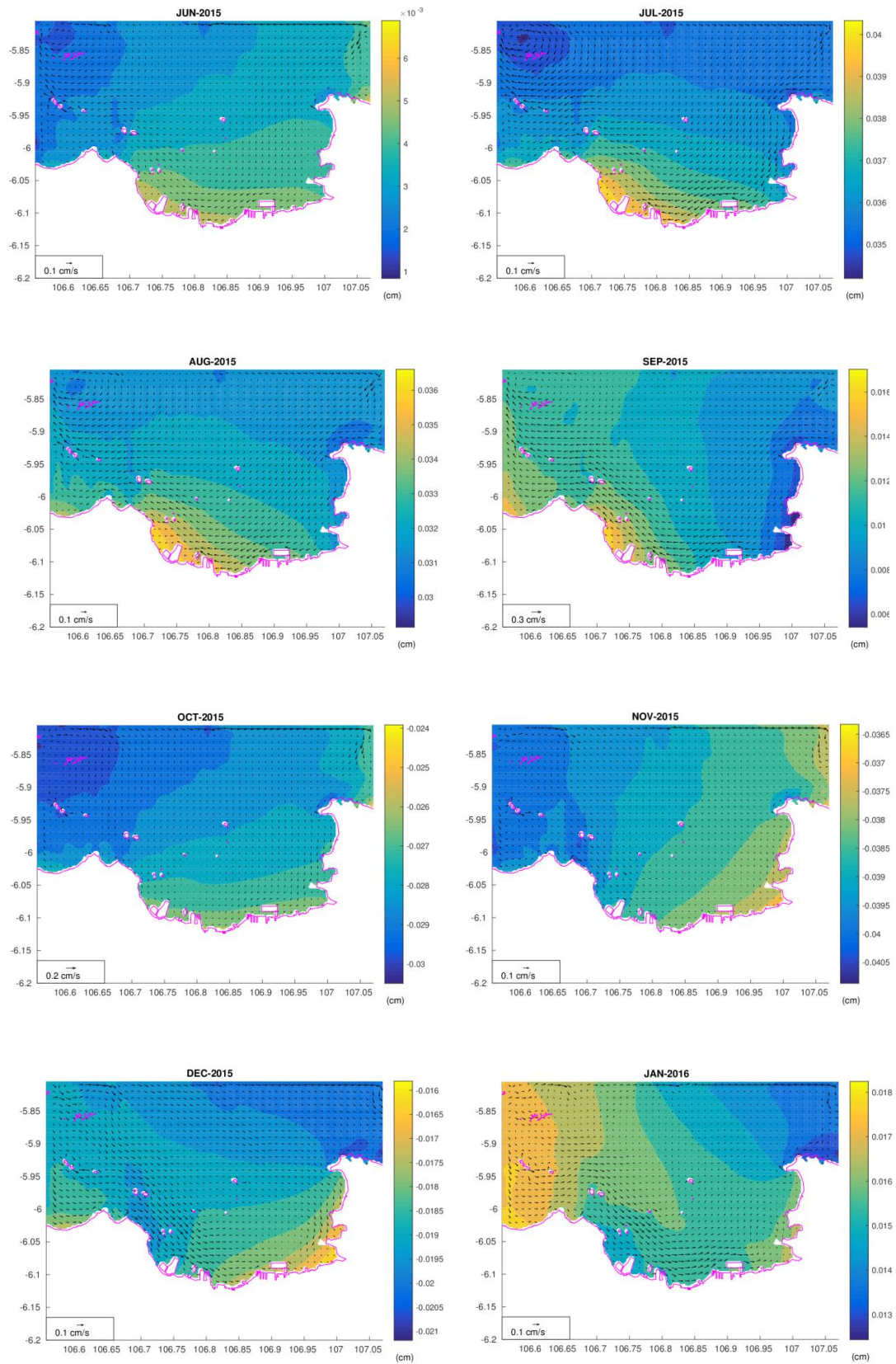
January 2016.

From March-May 2016 the surface velocity fields shows changing from southward to westward. This condition is similar to the propagation of high sea surfaces (the transition monsoon period). The easterly monsoon wind looks dominant in period May to August 2016. In this period the surface current propagate from east to west side of Jakarta Bay (westward currents). This situation confirmed to the existence of high sea surface in west side of Jakarta Bay.

The westward surface current fields increase from September to December 2016. This situation is also coinciding to the presence of high sea surface in east side of Jakarta Bay.

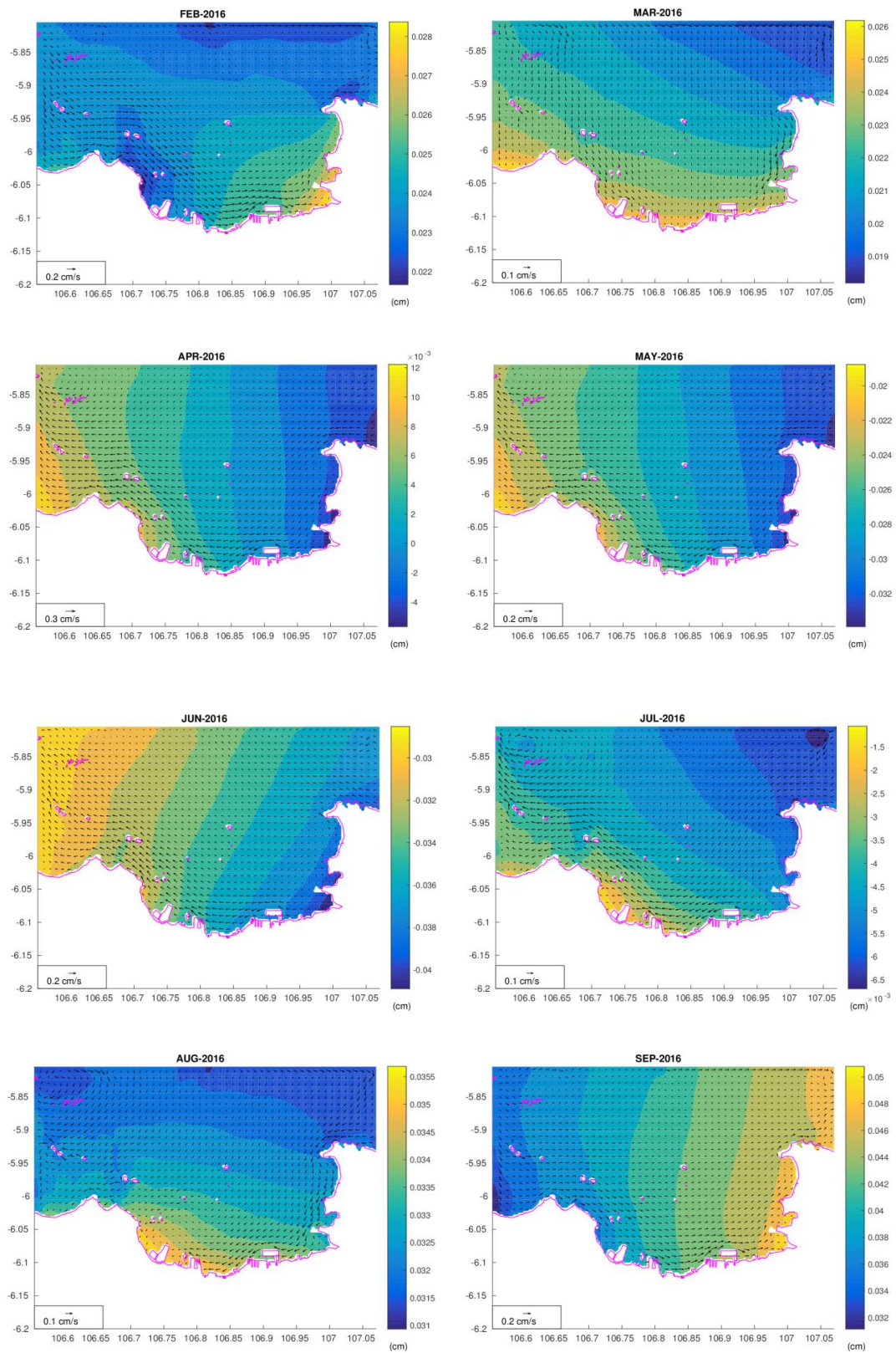
### **Perspective in the future**

To complete this physical-ecosystem model in Jakarta bay, we still find the good river discharges data surrounding the bay, at least between June 2015 to 2016.

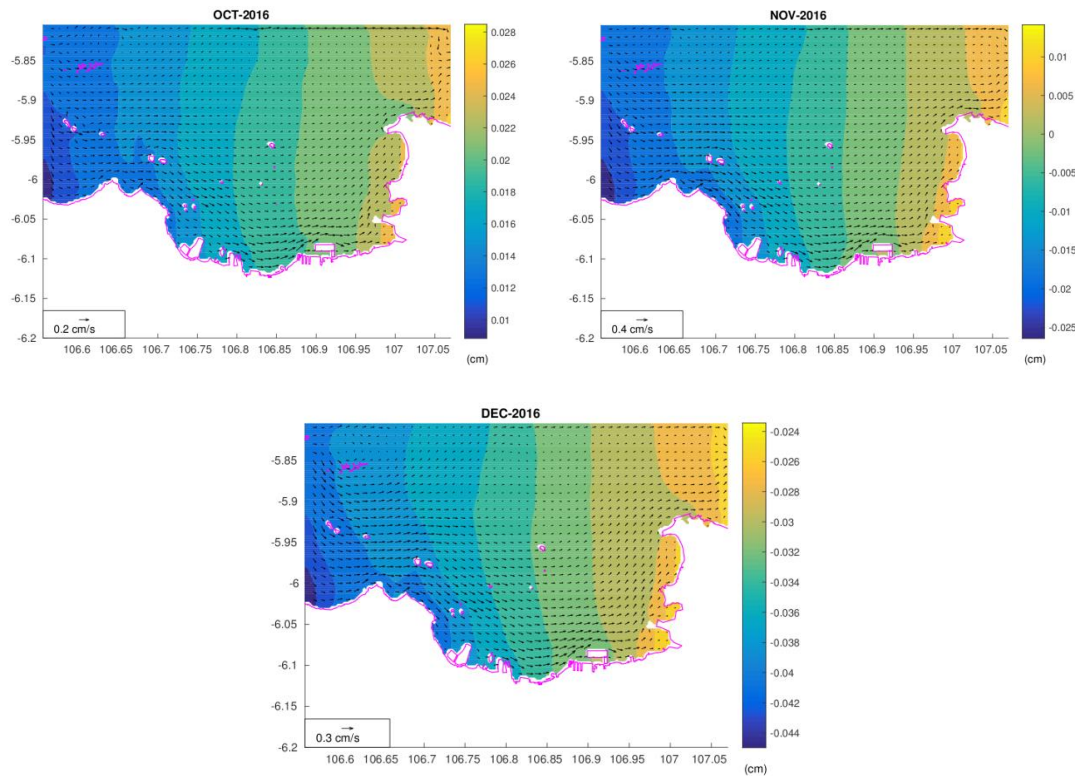


**Figure-1.** Monthly mean of Surface Current and Sealevel from June 2015 to December 2016  
(to be continued)





**Figure-1.** Monthly mean of Surface Current and Sealevel from June 2015 to December 2016  
(to be continued)



**Figure-1.** Monthly mean of Surface Current and Sealevel from June 2015 to December 2016

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